

Soil Biology & Biochemistry 37 (2005) 93-100

Soil Biology & Biochemistry

www.elsevier.com/locate/soilbio

Simulation of nitrogen rhizodeposition and assimilation back into corn (Zea mays L.) roots

J.A.E. Molina^{a,*}, C.E. Clapp^b, R.R. Allmaras^b, M.F. Layese^a

^aDepartment of Soil Water and Climate, University of Minnesota, St Paul, MN 55108, USA ^bUSDA-ARS and Department of Soil, Water and Climate, University of Minnesota, St Paul, MN 55108, USA

Received 10 November 2003; received in revised form 28 May 2004; accepted 25 June 2004

Abstract

The objectives were to quantify the dynamics of N exchange between corn root and soil under field conditions and to compute the amount of N-rhizodeposition that is recycled back into corn by maturity. In a previous publication the simulation model NCSWAP/NCSOIL was used to quantify the release of organic-C from corn and its incorporation into soil organic matter. In this publication, results from the same simulations but pertinent to N are presented. The model was calibrated against measured N concentration in corn and soil and ¹⁵N enrichment data obtained from a long-term field experiment located near St Paul, MN. Field management included rototillage, the removal of stover-residue and grain and 4 fertilizer N treatments: 200 and 20 Kg N ha⁻¹ added yearly from 1980 to 1992, with ¹⁵N added with the fertilizer from 1980 until either 1985 or 1992. The same total amount of ¹⁵N was added with the 200 and 20 Kg N ha⁻¹ fertilizer: 4 and 40 ¹⁵N at.%, respectively. Corn roots released 24% of total N uptake. This loss was mitigated by the recycling of 14% of N-rhizodeposition into corn by maturity. ¹⁵N enrichment in corn and soil was higher for the 200 than the 20 Kg N ha⁻¹ treatment. This resulted from the rapid N mineralization–immobilization turnover that channelled N through the inorganic N pool whose ¹⁵N enrichment was fixed yearly to that of the fertilizer. Tracer N enrichment decreased more rapidly in corn than in soil from 1986 to 1992 when tracer N was no longer added with the fertilizer, and by 1992, ¹⁵N was localized in the stable pool and flushed from the more labile pools.

© 2004 Elsevier Ltd. All rights reserved.

Keywords: Rhizosphere; Mineralization-immobilization turnover; Exudates; Carbon cycle; N fertilization; NCSWAP; NCSOIL

1. Introduction

Rhizodeposition of C from plant roots is well documented. Photosynthates are transferred below ground, leaked as exudates or sloughed off rootlets and decayed in the rhizosphere (Balesdent and Balabane, 1996; Clapp et al., 2000; Molina et al., 2001). Along with C, N is also released, incorporated into the soil organic matter (SOM) and made available to plants as NH₄⁺ by the mineralization–immobilization turnover (MIT). For example, up to 57 and 49% of N assimilated by legumes and barley, respectively, were found in the soil after removal of the roots at maturity (Jansen, 1996a). The decay of rhizodepositions from N₂ fixing plants

E-mail address: jamolina@umn.edu (J.A.E. Molina).

released NH₄⁺ which was available to plants growing in association with legumes (Jansen, 1996b). N released from alfalfa and birdsfoot trefoil originated more through root and nodules decay than exudations (Duback and Ruselle, 1994). N exchange was observed between the roots of an alfalfa-bromegrass sward (Walley et al., 1996).

The high C:N ratio of rhizodeposition from non-leguminous plants induces N immobilization: N is captured in the SOM with a net N loss for the plant from which it originated, even if this N is made available later on for subsequent crops (Huggins and Pan, 1993). This model of N dynamics in the rhizosphere assumes that the rate of net N mineralization from SOM is so slow as to preclude significant recycling of N to the source plant before maturity. Recent publications, however, would indicate that such is not the case. Both the amount of N released by roots and its rate of turnover are high. Høgh-Jensen and Schjoerring (2001) showed that N-rhizodeposition from

^{*} Corresponding author. Address: Department of Soil Science, University of Minnesota, 1991 Upper Buford Circle, 439 Borlaug Hall, St Paul, MN 55108, USA. Tel.: +1-612-625-6259; fax: +1-612-625-2208.

ryegrass exceeded the amount of N present in stubble. The recycling of root-derived N back into winter wheat during the growing season was demonstrated by Jimenez et al. (2002). They estimated that up to 54% of the N assimilated by winter wheat was leaked as rhizodeposition, a loss which was, however, mitigated by the existence of a 'N loop between soil and growing plants'.

Carbon and nitrogen dynamics in the soil-root system are linked. In a previous publication (Molina et al., 2001) we used the simulation model NCSWAP/NCSOIL to quantify the release of organic C from corn and its incorporation in the SOM. The model, calibrated against field data from a long-term experiment with continuous corn, simulated the kinetics of corn above-ground production, SOM, soil δ ¹³C values, and SOM derived from corn. In this publication, results from the same simulations but pertinent to N are presented. Our objectives were to quantify the dynamics of N exchange between corn roots and soil under field conditions and to compute the amount of N-rhizodeposition that is recycled back into corn by maturity.

2. Materials and methods

2.1. Field experiments and measured data

Data were obtained from ^{15}N microplots (1×3 m) located in the center of plots $(6 \times 9 \text{ m})$ established adjacent to the nitrogen, tillage, residue management (NTRM) longterm experiment started in 1980 at the University of Minnesota Outreach, Research and Education Park, Rosemount, MN. The 15N microplots were surrounded by a 30 cm metal barrier pushed 20 cm into the soil. Both field experiments are on-going long-term continuous corn studies. The ¹⁵N microplot experiment is a factorial of two rates of N application (200 and 20 kg N ha⁻¹), two corn stover-residue managements (stover removed or returned), and two methods of tillage (garden-size rotary tilled or not tilled). Detailed accounts of the ¹⁵N microplots are given by Clay et al. (1985, 1989) and Gollany et al. (2004), whereas details of the adjacent NTRM experiment is detailed by Clapp et al. (2000). This publication presents the data obtained from 1980 to 1992 for the following treatments of the ¹⁵N microplots: soil tilled; stover-residue plus grain removed; ammonium sulfate fertilizer N surface applied every year in spring at 200 and 20 kg N ha⁻¹ with 4 and 40 ¹⁵N at.%, respectively, for a total addition of about $8.0 \text{ kg}^{15} \text{N ha}^{-1}$. The $^{15} \text{N at.}\%$ of the added fertilizer N was measured every year. In 1985 the ¹⁵N microplots were split in two by a metal barrier similar to the one surrounding the plot. One half continued to receive ¹⁵N but the other half did not in order to document the relaxation kinetics of ¹⁵N disappearance from plant and soil. There were thus 4 N treatments (2 rates of N fertilizer application ×2 number of years when tagged N fertilizer was added). The soil at the site is a Waukegan silt loam (fine-silty over sandy or

sandy-skeletal, mixed, mesic Typic Hapludoll). Soil depth to gravel is about 100 cm. The climate is continental (average temperatures in January and July are -12 and 22 °C, respectively, with an annual average precipitation of 840 mm). The ¹⁵N plots were rototilled in the fall to a depth of 15 cm or less. Soil samples were taken from 0 to 15 and 15 to 30-cm layers, sieved (2 mm), and analyzed for total and tracer N.

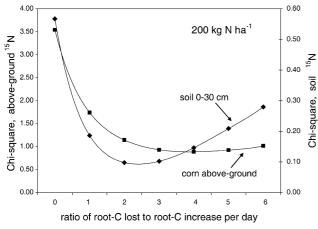
Nitrogen from soil and above-ground corn samples was determined using an elemental analyzer interfaced with a stable isotope ratio mass spectrometer (Carlo Erba, model NA1500 and Fisons, model Optima, Middlewich-Cheshire, UK). Specific field management, soil sampling and analytical techniques were detailed in Clapp et al. (2000).

2.2. Simulated data

Simulated N data were computed from the model NCSWAP/NCSOIL (Molina et al., 1983; Hadas and Molina, 1993; Molina et al., 1997). The source code is available at the web site http://www.soils.umn.edu/research/ ncswap-ncsoil/. The model simulates C and N-total and tracer-transformations related to soil water, thermal and solute transport in a soil profile. Optimum above-ground plant growth is represented by functions parameterized for mass increase and N% decrease measured for a reference year when N and water stress are minimal. Plant growth is adjusted for water and N stress by reduction factors. Adjustment to the temperature is achieved by linking the actual and reference air temperatures through degree day accumulation. Root and above ground daily mass increase are linked by a time dependent shoot-to-root ratio. Repartition of root mass in soil layers is controlled by the fraction of the daily root mass increase that is added to the top layers of the profile; the repartition is defined based upon physiological stages of the crop. The root mass not included in the top layers is distributed in the subjacent layers; roots grow until the beginning of decay at harvest or maturity; rhizodeposition decays as soon as formed. The amount of rhizodeposition was specified as the ratio of root-C lost as rhizodeposit to root-C increase per day. The ratio was obtained by calibration of simulated against experimental relaxation ¹⁵N kinetics.

N and C transformations are simulated by NCSOIL as a subroutine of NCSWAP. SOM in each 5-cm layer in the 100-cm profile is represented by four pools: labile and recalcitrant Pool I, Pool II, and Pool III. Organic debris in each layer is: (1) roots and (2) rhizodeposition from current year; (3) roots and (4) rhizodeposition accumulated from previous years; (5–7) three different above-ground crop residues; and (8–11) four different organic amendments.

Simulated rates of change in tracer N were computed at each time step as the rate of change in total N multiplied by the ratio of tracer to total N. No discrimination between tracer and total N rates of transformation was assumed. Initial tracer N concentrations of the soil pools were set to



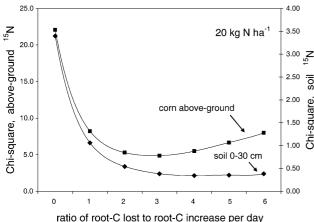


Fig. 1. Sensitivity of ^{15}N excess at.% in above-ground corn and soil (0–30 cm) to variations in the rate of rhizodepositions at 2 fertilizer treatments: 200 and 20 kg N ha $^{-1}$.

zero. Tracer N enrichments were expressed as excess ¹⁵N at.% relative to the measured average natural tracer N concentration of the soil, 0.37675 ¹⁵N at.%. Gollany et al. (2004) also give details of the simulation procedure using both ¹⁵N and ¹³C tracer measurements.

The simulated time course of N released from the roots was obtained by tagging N-rhizodeposition. For purposes of this research all pathways of C and N losses from the roots (exudation, passive leakage, tissue decomposition) were treated similarly as rhizodeposition. As N was lost from the root for the first time, the ratio of tracer to total N-rhizodeposition was set to 1.0. For subsequent time steps, the ratio was maintained to 1.0 after accounting for the tracer N recycled through the SOM. This simulation analysis was carried out for the year 1980 with nontagged 200 kg N ha⁻¹ fertilizer. The model assumed that the rapid microbial decay of N-rhizodeposition precluded any direct assimilation of organic N by the roots (Owen and Jones, 2001). The parameters that control C and N transformations in soil—such as half-life of the SOM pools, microbial efficiency of C assimilation and humification—were assigned the same values as used in previous simulations.

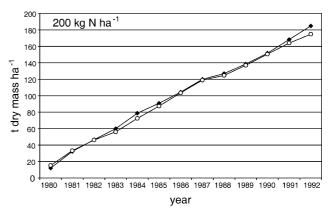
3. Results

3.1. Calibration

The rate of rhizodeposition was obtained for the value corresponding to a minimum of the χ^2 -function taken for measured and simulated above-ground and soil (0–30 cm) $^{15}{\rm N}$ relaxation kinetics at 200 and 20 kg N ha $^{-1}$ in 1980 (Fig. 1). The χ^2 -values dropped sharply from 0 to 2 C released for each C root produced per day and then changed little as the rate increased. Consequently, the daily ratio of the released C to root-produced C was set to 3 for later simulations. A minimum of the χ^2 -function could not be obtained for variations in the N% of the rhizodepositions which was set to 0.4 that of the above-ground corn.

3.2. Dry mass and nitrogen

The measured cumulative above-ground dry masses for the high and low N treatments were similar from 1980 to 1983. Subsequently, they gradually diverged (Fig. 2). This trend was reflected in the simulated data. Simulated N concentration in the above-ground corn was diluted from 4.9% at emergence to about 1% at harvest (Fig. 3). The measured N% at harvest was lower for the low N than for



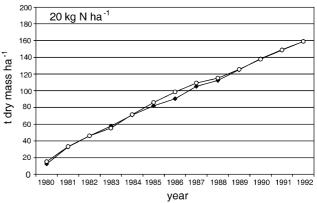
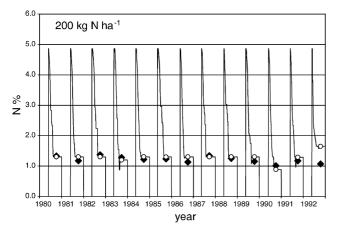


Fig. 2. Cumulative above-ground dry mass for the 200 and 20 kg N ha⁻¹ fertilizer treatments: (\bigcirc) simulated at time corresponding to measured data, (\blacklozenge) measured.



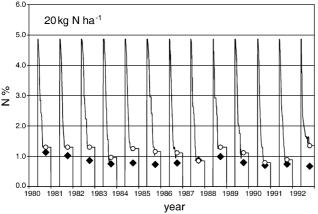


Fig. 3. Above-ground N percentage for the 200 and 20 kg N ha⁻¹ fertilizer treatments: (\bigcirc) simulated, (\bigcirc) simulated at time corresponding to measured data, (\blacklozenge) measured.

the high N treatment, even for the first 4 years when there was no distinction between the dry mass accumulation at 20 and 200 kg N ha $^{-1}$. A lower N% with the low N treatment was also simulated but not sufficiently depressed to closely match the measured data (Fig. 3).

The simulated inorganic plus organic N (plant root residues excluded by sieving) decreased gradually at the same rate in the 0–15 and 15–30 cm soil layers (Fig. 4). The addition of 200 kg N ha^{-1} was marked in the top layer by simulated peaks that disappeared with plant N uptake and nitrate leaching. Measured and simulated data were matched in the top layer for both N rates. In the 15–30 cm layer, however, the measured data did not fit the simulated trends and were highly variable in the high N treatment.

3.3. Tracer nitrogen

The measured and simulated dynamics of ¹⁵N in the above-ground portion of corn and in soil are shown in Figs. 5–8. Tracer N enrichment in corn and soil both differed with the ¹⁵N concentration of the fertilizer N (40 and 4 ¹⁵N at.% for the 20 and 200 kg N ha⁻¹, respectively; notice the different excess ¹⁵N at.% scales) although the same

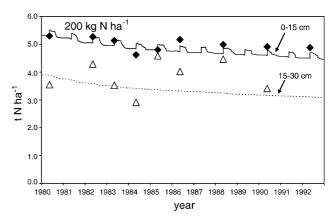
amount of tracer N was applied, i.e. $8.0 \text{ kg}^{15} \text{N ha}^{-1}$. ¹⁵N enrichment was higher in both corn and soil for the low fertilizer N application.

Every year, a sudden increase in ¹⁵N enrichment was simulated in the 0–15 cm soil layer corresponding to the addition of ¹⁵N enriched fertilizer N. As inorganic N subsided with corn N uptake and mass transport, this ¹⁵N peak stabilized to the ¹⁵N enrichment of the SOM. The peaks were less pronounced and occurred with a time lag in the 15–30 cm layer. A better fit between measured and simulated ¹⁵N enrichment was obtained in the 15–30 than in the 0–15 cm soil layers; a pattern which is the reverse of that observed with total soil N.

The measured relaxation kinetics exhibited in soil and plant from the microplots that did not receive ¹⁵N from 1985 to 1992 was accounted for by the simulation. Tracer N enrichment decreased more rapidly in plant than in soil. Tracer N in soil gradually stabilized at about one-half the level attained in 1985. ¹⁵N enrichment in corn decreased suddenly in 1985 to an almost constant low level in subsequent years.

3.4. N exchange between corn and soil

The simulated dynamic of rhizodeposition ¹⁵N for the year 1980 and in the high N treatment showed that corn



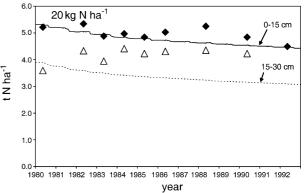
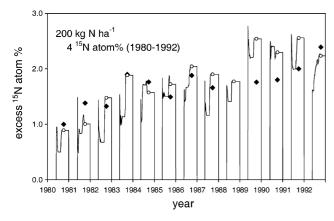


Fig. 4. N concentration in soil for the 200 and $20 \, \text{kg N ha}^{-1}$ fertilizer treatments: (\longrightarrow) simulated, (\spadesuit) measured 0–15 cm; (- - -) simulated, (\triangle) measured 15–30 cm.



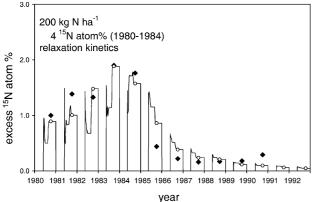


Fig. 5. Tracer ^{15}N in above-ground corn for the 200 kg N ha^{-1} with 4 ^{15}N at.%, and 2 periods of ^{15}N applications: 1980–1992 and 1980–1984: () simulated, () simulated at time corresponding to measured data, (•) measured.

roots had leaked 63 kg N ha⁻¹ or 24% of the N uptake at maturity (Fig. 9). This was not a total loss for corn because 9 kg N ha⁻¹ was reabsorbed by the roots for a net loss of 54 kg N ha⁻¹. Of this net loss, one-half was partitioned at maturity to the SOM, and 30% to inorganic N while 20% remained in the rhizodepositions.

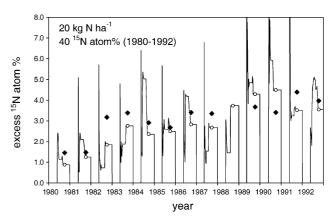
4. Discussion

The level of rhizodeposition was obtained by calibration against ¹⁵N enrichment kinetics. The same approach used with the C kinetics gave a wide range of possible values because this parameter had little effect on the simulated kinetics (Molina et al., 2001): most of the information was lost with the unmonitored escape of CO₂. With N, however, the information was kept in the system as losses by denitrification and mass flow were not as extensive as those associated with CO₂.

The trend of measured kinetics for total and tracer N enrichment in corn and the two soil layers was reproduced by the simulation. The ¹⁵N concentration effect and the relaxation curves were accounted for by the model. The simulated kinetics resulted from several processes that run simultaneously, competed for N and were quantitatively

integrated by the model. Inorganic N is assimilated by corn and immobilized by microbes; the N mineralization—immobilization turnover (MIT; Jansson and Persson, 1982) driven by the microbial activity (microbial succession; Beek and Frissel, 1973; Molina and Smith, 1988) shuttles N between NH₄⁺ and the microbial biomass. The MIT also stabilizes some N in the SOM which decays at a lower rate than the microbial population, thus gradually reducing the rate of MIT and N accessibility to corn. Rhizodepositions recycle N in soil. They also provide the energy and C for N immobilization and microbial growth. The former process reduces N accessibility, and the latter increases it by stimulating the rates of rhizodeposition decay and MIT.

The model uses four SOM pools: two for the microbial biomass (Pool I, labile and resistant) and two more stabilized pools (Pool II and III). A very rapid rate of microbial decay was required to obtain the simulated kinetics (2 and 17 d for the half life of the labile and resistant microbial pools). These values were used throughout the simulations carried out with the models NCSWAP/NCSOIL and NCSOIL (e.g. Nicolardot et al., 1994). They are also in the order of the decay rates measured for live or dead microbes as well as microbial components added to soil (Nelson et al., 1979). It is this



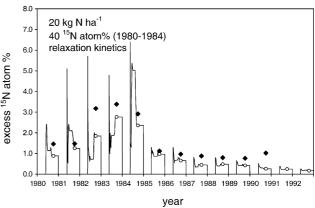
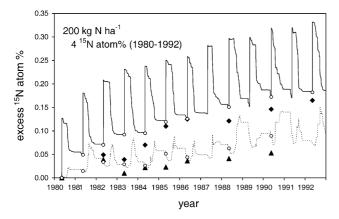


Fig. 6. Tracer ¹⁵N in above-ground corn for the 20 kg N ha⁻¹ with 40 ¹⁵N at.%, and 2 periods of ¹⁵N applications: 1980–1992 and 1980–1984: (→) simulated, (○) simulated at time corresponding to measured data, (◆) measured.



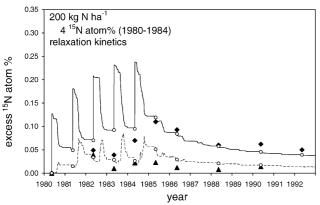


Fig. 7. Inorganic plus organic tracer 15 N in soil for the 200 kg N ha $^{-1}$ with 4 15 N at.%, and 2 periods of 15 N applications: 1980–1992 and 1980–1984: simulated (\bigcirc 0–15 cm, (- - -) 15–30 cm; (\bigcirc) simulated at time corresponding to measured data; measured (\spadesuit) 0–15 cm, (\blacktriangle) 15–30 cm.

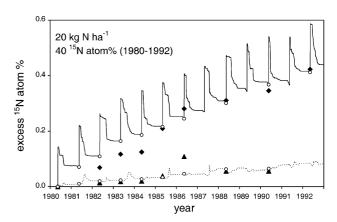
short half-life which drives the MIT at a rate fast enough to recycle 14% of the N-rhizodeposition back into the corn before it reaches maturity. The ¹⁵N enrichment of the inorganic N recycled in corn was different from that of the rhizodeposition from which it was derived as it was diluted in the NH₄⁴ pool during the MIT.

The simulated gross mineralization and immobilization in 1980 were 373 and $319 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ in the top 0-15 cm soil layer, respectively. It corresponds to an average 1.57 and 1.35 µg N g⁻¹ d⁻¹, respectively; these values are in the order of those found for a soil incubated under optimum laboratory conditions (Broadbent, 1965; Molina et al., 1990). Gross N mineralization rates ranging from 0.21 to 4.66 μ g N g⁻¹ d⁻¹ were reported by Watkins and Barraclough (1996), and Monaghan and Barraclough (1997). For the whole profile and over one year, these rates translated to very high levels of MIT. In 1980, the simulated gross N mineralization and immobilization were 3577 and 3017 kg N ha⁻¹, respectively. It provided 559 kg mineralized $NH_4^+ - N ha^{-1}$, a high value consistent with a soil fertility level able to sustain maximum corn yield for 4 years with only 20 kg fertilizer N ha⁻¹.

Although the same amount of 15 N (8.0 kg 15 N ha $^{-1}$) was added with the two levels of fertilizer N (200 and 20 kg N ha^{-1}), the different enrichments

(40 and 4 15N at.%) induced differences in the soil and plant ¹⁵N enrichment. This concentration effect was caused by the obligatory passage of ¹⁵N though the inorganic N pool. The intense rate of MIT which shuttled N between the organic and inorganic pools caused the adjustment of the organic ¹⁵N enrichment to that of the inorganic pool fixed once a year by the fertilizer addition. Inorganic N losses by denitrification and leaching in this 15N experiment were reported by Gollany et al. (2004). When corn stover was removed, the simulated denitrification losses were less and NO₃ leaching losses were greater than when crop residue was returned. Simulated losses by both means were greater at 200 compared to 20 kg N ha⁻¹ fertilization. Predicted losses of N were no more than 25% of applied N, but were somewhat larger than estimates obtained from most field observations (Gollany et al., 2004).

When ¹⁵N was not added with the fertilizer N from 1985 to 1992, the sudden fall of ¹⁵N enrichment in corn may have originated from a reduced N availability from soil, the fall of the SOM ¹⁵N enrichment, or a combination of those factors (Figs. 5 and 6). By contrast, the soil ¹⁵N enrichment decrease was more gradual, and leveled to about one-half of its peak values attained in 1984 (Figs. 7 and 8). These features could indicate that different SOM pools were exploited by the plant roots. Assuming that plant N



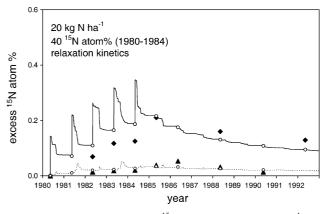


Fig. 8. Inorganic plus organic tracer 15 N in soil for the 20 kg N ha $^{-1}$ with 40 15 N at.%, and 2 periods of 15 N applications: 1980–1992 and 1980–1984: simulated (----) 0–15 cm, (---) 15–30 cm; (\bigcirc) simulated at time corresponding to measured data; measured (\spadesuit) 0–15 cm, (\spadesuit) 15–30 cm.

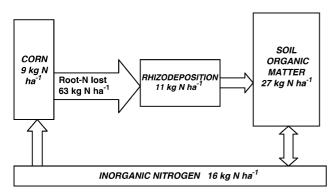
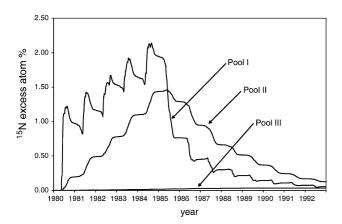


Fig. 9. Flow from emergence to maturity and accumulation at maturity of corn N rhizodeposition.

absorption was achieved through inorganic N exclusively, the change in ¹⁵N enrichment in corn and soil would be controlled by the decay rates of the various SOM pools. The simulated relaxation kinetics were obtained concomitant to a gradual disappearance of ¹⁵N content in Pool I and Pool II, and accumulation in the stable Pool III (Fig. 10). By 1982, Pool I and Pool II had lost almost all the added ¹⁵N (0.044 and 0.162 kg ¹⁵N ha⁻¹, respectively), whereas Pool III had accumulated 1.49 kg ¹⁵N ha⁻¹ (0.755 µg ¹⁵N g⁻¹ soil) in the 0–15 cm layer. By contrast, the soil pools could not be distinguished in terms of ¹⁵N enrichment (Fig. 10). These



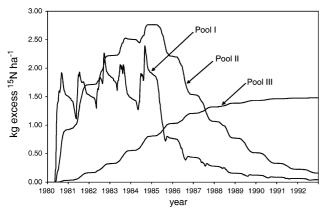


Fig. 10. 15 N enrichment and concentration in the SOM pools in the 0–15 cm layer, for the 200 kg N ha $^{-1}$ treatment with 15 N addition discontinued in 1985.

circumstances could provide a key to experimentally differentiate Pool III from the other pools: Chemical fractions from 0 to 15 cm layer of 1992 soil samples should have a low ¹⁵N enrichment but high ¹⁵N content to represent Pool III

References

Balesdent, J., Balabane, M., 1996. Major contribution of roots to soil carbon storage inferred from maize cultivated soils. Soil Biology & Biochemistry 28, 1261–1263.

Beek, J., Frissel, M., 1973. Simulation of nitrogen behaviour in soils. Pudoc, Wageningen, The Netherlands.

Broadbent, F.E., 1965. Effect of fertilizer nitrogen on the release of soil nitrogen. Soil Science Society of America Proceedings 29, 692–696.

Clapp, C.E., Allmaras, R.R., Layese, M.F., Linden, D.R., Dowdy, R.H., 2000. Soil organic carbon and ¹³C abundance as related to tillage, crop residue, and nitrogen fertilization under continuous corn management in Minnesota. Soil and Tillage Research 55, 127–142.

Clay, D.E., Clapp, C.E., Molina, J.A.E., Linden, D.R., 1985. Nitrogentillage-residue management: 1. Simulating soil and plant behavior by the model NCSWAP. Plant and Soil 84, 67–77.

Clay, D.E., Clapp, C.E., Linden, D.R., Molina, J.A.E., 1989. Nitrogentillage-residue management: 3. Observed and interactions among soil depth, nitrogen mineralization, and corn yield. Soil Science 147, 319–327.

Duback, M., Ruselle, M., 1994. Forage legume roots and nodules and their role in nitrogen transfer. Agronomy Journal 86, 259–266.

Gollany, H.T., Molina, J.A.E., Clapp, C.E., Allmaras, R.R., Layese, M.F., Baker, J.M., Cheng, H.H., 2004. Denitrification and nitrogen leaching in continuous corn as related to residue and nitrogen management. Environmental Management 33, 5289–5298.

Hadas, A., Molina, J.A.E., 1993. Simulation of nitrogen assimilation by heterotrophic soil microbial biomass. Physiology Plantarum 87, 528–534.

Høgh-Jensen, H., Schjoerring, J.K., 2001. Rhizodeposition of nitrogen by red clover, white clover and ryegrass leys. Soil Biology & Biochemistry 33, 439–448.

Huggins, D.R., Pan, W.L., 1993. Nitrogen efficiency component analysis: an evaluation of cropping system differences in productivity. Agronomy Journal 85, 898–905.

Jansen, E.S., 1996a. Rhizodeposition of N by pea and barley and its effect on soil N dynamics. Soil Biology & Biochemistry 28, 65–71.

Jansen, E.S., 1996b. Barley uptake of N deposited in the rhizosphere of associated pea. Soil Biology & Biochemistry 28, 159–168.

Jansson, S.L., Persson, J., 1982. Mineralization and immobilization of soil nitrogen, In: Stevenson, F.J., Bremner, J.M., Hauck, R.D., Keeney, D.R. (Eds.), Nitrogen in Agricultural Soils, vol. 22. Agronomy Monograph, Madison, WI, pp. 229–252.

Jimenez, M.A., Schmid, H., von Lützow, M., Guster, R., Munch, J.C., 2002. Evidence for recycling of N from plants to soil during the growing season. Geoderma 105, 223–241.

Molina, J.A.E., Smith, P., 1988. Modeling carbon and nitrogen processes in soils. Advances in Agronomy 62, 253–298.

Molina, J.A.E., Clapp, C.E., Shaffer, M.J., Chichester, F.W., Larson, W.E., 1983. NCSOIL, a model of nitrogen transformations in soil: description, calibration and behavior. Soil Science Society of America Journal 47, 85–91.

Molina, J.A.E., Hadas, A., Clapp, C.E., 1990. Computer simulation of nitrogen turnover in soil and priming effect. Soil Biology & Biochemistry 22, 349–353.

Molina, J.A.E., Crocker, G.J., Grace, P.R., Klir, J., Korschens, M., Poulton, P.R., Richter, D.D., 1997. Simulating trends in soil organic

- carbon in long-term experiments using the NCSOIL and NCSWAP models. Geoderma 81, 91–107.
- Molina, J.A.E., Clapp, C.E., Linden, D.R., Allmaras, R.R., Layese, M.F., Dowdy, R.H., Cheng, H.H., 2001. Modeling the incorporation of corn (*Zea mays* L.) carbon from roots and rhizodeposition into soil organic matter. Soil Biology & Biochemistry 33, 83–92.
- Monaghan, R., Barraclough, D., 1997. Contribution to N mineralization from soil macroorganic matter fractions incorporated into two field soils. Soil Biology & Biochemistry 29, 1215–1223.
- Nelson, D.W., Martin, J.P., Ervin, J.O., 1979. Decomposition of microbial cells and components in soil and their stabilization through complexing with model humic acid-type phenolic polymers. Soil Science Society of America Journal 43, 84–88.
- Nicolardot, B., Molina, J.A.E., Allard, M.R., 1994. C and N fluxes between pools of soil organic matter: Model calibration with long-term incubation data. Soil Biology & Biochemistry 26, 235–243.
- Owen, A.G., Jones, D.L., 2001. Competition for amino acids between wheat roots and rhizosphere microorganisms and the role of amino acids in plant N acquisition. Soil Biology & Biochemistry 33, 651–657.
- Walley, F.L., Tomm, G.O., Matus, A., Slikard, A.E., Van-Kessel, C., 1996.
 Allocation and cycling of nitrogen in an alfalfa-bromegrass sward.
 Agronomy Journal 88, 834–843.
- Watkins, N., Barraclough, D., 1996. Gross rates of N mineralization associated with the decomposition of plant residues. Soil Biology & Biochemistry 28, 169–175.